



Hydrogen Escape from Mars is Driven by Seasonal and Dust Storm Transport of Water

<u>Shane W. Stone</u>,¹ Roger V. Yelle,¹ Mehdi Benna,² Daniel Lo,¹ Meredith K. Elrod,² Paul R. Mahaffy²

> ¹Lunar and Planetary Laboratory, University of Arizona ²NASA Goddard Space Flight Center

NTRODUCTION

- H escape to space could account for the loss of 85% of the initial water inventory of Mars over the last 4 billion years.
- Previous investigations have found seasonal variations of an order of magnitude or more in the upper atmospheric H density, implying a similar variation in the H escape rate. *chaffin et al.* (2014), *clarke et al.* (2014), *Bhattacharyya et al.* (2015), *Clarke et al.* (2017)
 - These variations are too rapid to be explained by the slow and steady delivery of H₂ from the lower atmosphere by diffusion. This is the "classical" source of H.
- Rapid seasonal and dust-storm-induced changes in the vertical distribution of H₂O have now been observed by MCS, SPICAM, ACS, and NOMAD. *Heavens et al.* (2018), *Fedorova et al.* (2018, 2019), *Vandaele et al.* (2019)
- H₂O transported to the upper atmosphere is rapidly destroyed and the H produced can escape efficiently. MAVEN Neutral Gas and Ion Mass Spectrometer (NGIMS) is uniquely positioned to collect *in situ* measurements of neutral and ionic species in the upper atmosphere.

CLASSICAL SOURCE OF H New Source of H 4) Jeans escape of H 4) Jeans escape of H **Upper Atmosphere Upper Atmosphere** CO_2^+ + H₂O \rightarrow H₂O⁺ + CO₂ $HCO^+ + H_2O \rightarrow H_2O^+ + CO$ $CO_{2}^{+} + H_{2} \rightarrow HCO_{2}^{+} + H$ 3) 3) $H_{2}O^{+} + e^{-} \rightarrow H_{2}O + H$ **Weak Hygropause** Hygropause 2) Diffusion of H₂ 2) Transport of H₂O $H_2O(g) \rightarrow H_2O(l)$ Lower Atmosphere Lower Atmosphere $H_2O(s) \rightarrow H_2O(g)$ $H + HO_{2} \rightarrow H_{2} + O_{2}$ 1) 1)

Shane Stone (stone@lpl.arizona.edu)

DIRECT TRANSPORT OF WATER

• Diffusion of H₂ is slow and steady, too slow to explain observed rapid order-of-magnitude variations in the exospheric H abundance.

- Transport of H₂O from the lower atmosphere, however, is fast enough. Fedorova et al. (2018, 2019), Heavens et al. (2018), Vandaele et al. (2019)
- NGIMS data can be used to investigate the diurnal, seasonal, and dust-related variation of water and water-related ions in the upper atmosphere of Mars.
 - We measure abundances of H_2O^+ and H_3O^+ .
 - We calculate H₂O abundances from NGIMS measurements assuming photochemical equilibrium and use direct measurements of H₂ from the NGIMS neutral mode.
- This allows us to differentiate between the transport of H₂O and H₂ to the upper atmosphere.

VARIATION OF WATER



The sinusoidal seasonal variation is apparent in the H₂O mixing ratio.

Dust storms lead to a significant increase in the H₂O mixing ratio over a short time period.

 There are no strong seasonal or dust-storm-induced variations in measured H₂ abundances.

VARIATION OF WATER



 The upper atmosphere (~150 km) contains >1 ppm H₂O throughout the Martian year.

- The transport of water into the upper atmosphere is seasonal, peaking in southern summer.
 - The largest water vapor maximum in the lower atmosphere occurs during northern summer.
- Southern summer occurs near perihelion and this is the season of high dust activity.

VARIATION OF WATER IONS



- H₂O⁺ is a chemical intermediate that lies between an injection of H₂O from below and H escape from the top of the atmosphere.
- There are clear diurnal and seasonal variations in the relative abundance of H₂O⁺.
- We observe an abrupt, marked increase in H₂O⁺ immediately following the onset of dust activity during 3 dust storms.
- Similar, but smaller, variations are observed in H₃O⁺.

CALCULATING WATER DENSITIES

 Assuming photochemical equilibrium, we construct simple equations for the calculation of H₂O abundance from NGIMS ion and CO₂ measurements:

 $HCO^{+} + H_2O \xrightarrow{k_1} H_3O^{+} + CO$ $H_2O^{+} + H_2O \xrightarrow{k_2} H_3O^{+} + OH$

$$H_{3}O^{+} + e^{-} \xrightarrow{\alpha_{1}} OH + H + H$$
$$\xrightarrow{\alpha_{2}} H_{2}O + H$$
$$\xrightarrow{\alpha_{3}} OH + H_{2}$$
$$\xrightarrow{\alpha_{4}} O + H_{2} + H$$

$$[H_2O] = (\alpha_1 + \alpha_2 + \alpha_3 + \alpha_4) \frac{[H_3O^+][e^-]}{k_1[HCO^+] + k_2[H_2O^+]}$$

DUST EVENTS



- Looking at the MY 34 global storm in greater detail, we observe a...
 - 3.1x increase in $[\mathbf{H}_{2}\mathbf{0}^{+}]/[e^{-}]$
 - 2.5x increase in $[\mathbf{H}_{3}\mathbf{0}^{+}]/[e^{-}]$

...over ~2 days.

 We do not observe significant change in [O₂⁺]/[e⁻], indicating that the storm did not perturb the entire ionosphere, but only the abundances of these water-related ions.

DUST EVENTS



During the MY 34 global storm, we observe a 2.4x increase in X(H₂O) over ~2 days.

No change in the H₂ mixing ratio is observed, indicating that H₂O, not H₂ is responsible for the observed perturbations in the ionosphere.

H PRODUCTION FROM WATER



 1D photochemical models were constructed for low H₂O and high H₂O cases.

	Low H ₂ O	High H ₂ O
H ₂ O Mixing Ratio at 80 km (ppm)	2	430
Net H ₂ O Destruction (cm ⁻² s ⁻¹)	2.6×10 ⁷	1.6×10^{9}
H Production (H ₂ O) (cm ⁻² s ⁻¹)	5.0×10 ⁷	2.9×10 ⁹
H ₂ Destruction Rate (cm ⁻² s ⁻¹)	9.6×10 ⁷	9.6×10 ⁷

• H produced from H₂O in the ionosphere can escape efficiently since it is produced close to the exobase.

CONCLUSIONS

- We observe diurnal, seasonal, and dust-storm-induced variations in upper atmospheric H₂O⁺ and H₃O⁺ abundances using data from NGIMS onboard MAVEN.
- These variations are due to the upward transport of H₂O past the hygropause and into the middle and upper atmosphere.
- The upper atmosphere contains >1 ppm H₂O throughout the Martian year.
 - Dust storms rapidly increase the upper atmospheric H₂O abundance by up to a factor of 2 over a few sols.
- Escaping H atoms are produced from H₂O near the exobase *via* reactions with ions.
- The contribution of H₂O to H escape is likely comparable to or greater than that of H₂.
 - A global dust storm leads to more than a Martian year's worth of H production and escape in just 45 days.